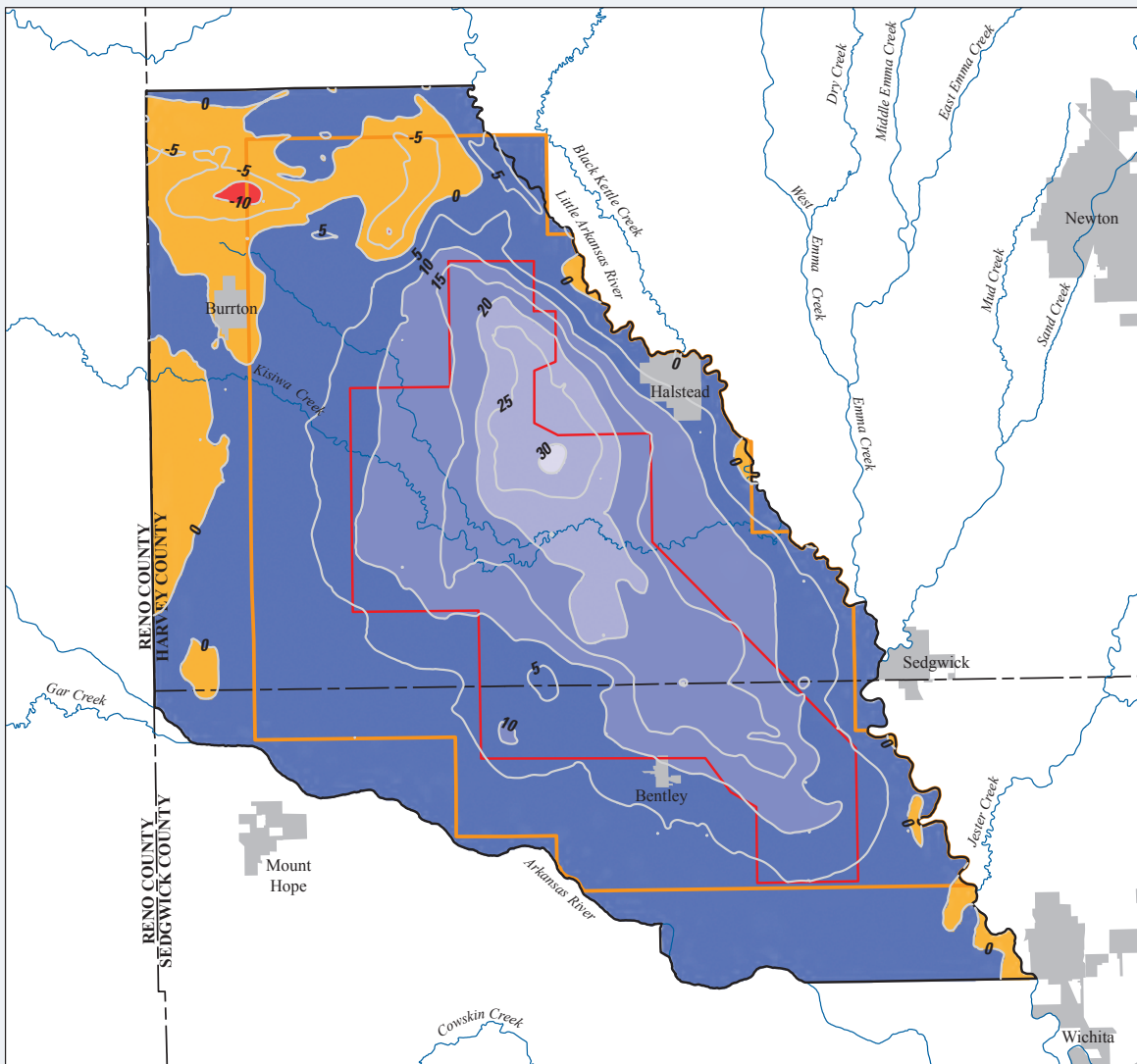


Prepared in cooperation with the City of Wichita, Kansas

Status of Groundwater Levels and Storage Volume in the *Equus* Beds Aquifer near Wichita, Kansas, January 2016



Scientific Investigations Report 2016–5165

Cover. Figure 8 from this report.

Status of Groundwater Levels and Storage Volume in the *Equus* Beds Aquifer near Wichita, Kansas, January 2016

By Brian J. Klager

Prepared in cooperation with the City of Wichita, Kansas

Scientific Investigations Report 2016–5165

**U.S. Department of the Interior
U.S. Geological Survey**

U.S. Department of the Interior
SALLY JEWELL, Secretary

U.S. Geological Survey
Suzette M. Kimball, Director

U.S. Geological Survey, Reston, Virginia: 2016

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Suggested citation:

Klager, B.J., 2016, Status of groundwater levels and storage volume in the *Equus* Beds aquifer near Wichita, Kansas, January 2016: U.S. Geological Survey Scientific Investigations Report 2016–5165, 15 p., <https://doi.org/10.3133/sir20165165>.

ISSN 2328-0328 (online)

Acknowledgments

The author acknowledges the assistance of staff of the city of Wichita and the staff of *Equus* Beds Groundwater Management District No. 2 for collecting and providing groundwater-level data used in this report.

Technical reviews provided by U.S. Geological Survey employees Virginia McGuire and Brian Kelly improved the technical and editorial clarity of this report.

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Conversion Factors

U.S. customary units to International System of Units

Multiply	By	To obtain
Length		
inch (in.)	2.54	centimeter (cm)
inch (in.)	25.4	millimeter (mm)
foot (ft)	0.3048	meter (m)
Area		
square mile (mi ²)	259.0	hectare (ha)
square mile (mi ²)	2.590	square kilometer (km ²)
acre	4,046.86	square meter (m ²)
acre	0.00404686	square kilometer (km ²)
acre	0.404686	hectare (ha)
Volume		
acre-foot (acre-ft)	1,233	cubic meter (m ³)
acre-foot (acre-ft)	0.325851	million gallons (Mgal)
Flow rate		
acre-foot per year (acre-ft/yr)	1,233	cubic meter per year (m ³ /yr)

Temperature in degrees Celsius (°C) may be converted to degrees Fahrenheit (°F) as follows:

$$^{\circ}\text{F}=(1.8\times^{\circ}\text{C})+32$$

Temperature in degrees Fahrenheit (°F) may be converted to degrees Celsius (°C) as follows:

$$^{\circ}\text{C}=(^{\circ}\text{F}-32)/1.8$$

Vertical coordinate information is referenced to the North American Vertical Datum of 1988 (NAVD 88).

Horizontal coordinate information is referenced to the North American Datum of 1983 (NAD 83).

Status of Groundwater Levels and Storage Volume in the *Equus* Beds Aquifer near Wichita, Kansas, January 2016

By Brian J. Klager

Abstract

The *Equus* Beds aquifer in south-central Kansas, which is part of the High Plains aquifer, serves as a source of water for municipal and agricultural users in the area. The city of Wichita has used the *Equus* Beds aquifer as one of its primary water sources since the 1940s. The aquifer in and around Wichita's well field reached historically low water levels in 1993, prompting the city to adopt new water-use and conservation strategies to ensure future water supply needs were met. Part of the plan was to initiate a managed aquifer recharge program called the *Equus* Beds Aquifer Storage and Recovery project. The goal of the managed aquifer recharge program is to artificially recharge the *Equus* Beds aquifer with treated water from the Little Arkansas River. As part of the *Equus* Beds Aquifer Storage and Recovery project, the city of Wichita and the U.S. Geological Survey have partnered in a long-term cooperative study to monitor and describe the quantity and quality of the water in the *Equus* Beds aquifer and the Little Arkansas River.

The city of Wichita, the *Equus* Beds Groundwater Management District No. 2, the Kansas Department of Agriculture–Division of Water Resources, and the U.S. Geological Survey collected groundwater levels in numerous wells screened in the *Equus* Beds aquifer in the area in and around Wichita's well field in January 2016. The measurements were used to interpolate potentiometric surfaces for shallow and deep parts of the aquifer in the study area. These potentiometric surfaces were compared with potentiometric surfaces from previous years to estimate changes in water levels and storage volume in the study area.

Groundwater levels were generally higher in January 2016 than they were in January 2015. On average, in January 2016, groundwater levels in the shallow part of the aquifer were about 3.4 feet higher and groundwater levels in the deep part of the aquifer were about 3.8 feet higher than in January 2015. The volume of water stored in the study area decreased by about 74,000 acre-feet between predevelopment (the time period before substantial pumpage began in the 1940s) and January 2016; increased by about 121,000 acre-feet between the historic low in 1993 and January 2016; and increased by about 61,000 acre-feet between January

2015 and January 2016. About 62 percent of the storage volume lost between predevelopment and 1993 has been recovered. The increase in storage volume from January 2015 to January 2016 can probably be attributed to less pumping by the city of Wichita and irrigators, more recharge due to higher-than-average precipitation, and higher volumes of artificial recharge in 2015.

Introduction

The *Equus* Beds aquifer, which is part of the High Plains aquifer, is an important source of water for municipal and agricultural users in south-central Kansas (*Equus* Beds Groundwater Management District No. 2, 2008). The U.S. Geological Survey (USGS) and the city of Wichita, Kansas, have been engaged in a long-term cooperative partnership to study the quantity and quality of water in the *Equus* Beds aquifer and the Little Arkansas River since city well field operations began in the 1940s. The central Wichita well field (WWF) (fig. 1) in the *Equus* Beds aquifer is one of the primary water-supply sources for the city of Wichita. The aquifer reached historically low water levels in 1993, prompting the city to adopt the Integrated Local Water Supply Plan (ILWSP) (Warren and others, 1995). The ILWSP included plans to increase use of surface water from Cheney Reservoir and increase water-conservation efforts, as well as a plan to help ensure adequate future water supply called the *Equus* Beds Aquifer Storage and Recovery (ASR) project. The goal of the ASR project is to artificially recharge the *Equus* Beds aquifer to increase aquifer storage and to allow for later recovery of the groundwater. This is accomplished by withdrawing water from the Little Arkansas River, treating the water using drinking-water quality standards as a guideline, and then injecting the water through wells or allowing it to infiltrate through recharge basins into the aquifer. Increases in storage within an area called the basin storage area (BSA) (fig. 1) attributable to the ASR project are credited to the city of Wichita by the State of Kansas, and those credits can be exchanged for the right to pump volumes of water from the aquifer greater than the city's normal annual water right. The increases in storage attributable to the ASR project are

2 Status of Groundwater Levels and Storage Volume in the *Equus* Beds Aquifer near Wichita, Kansas, January 2016

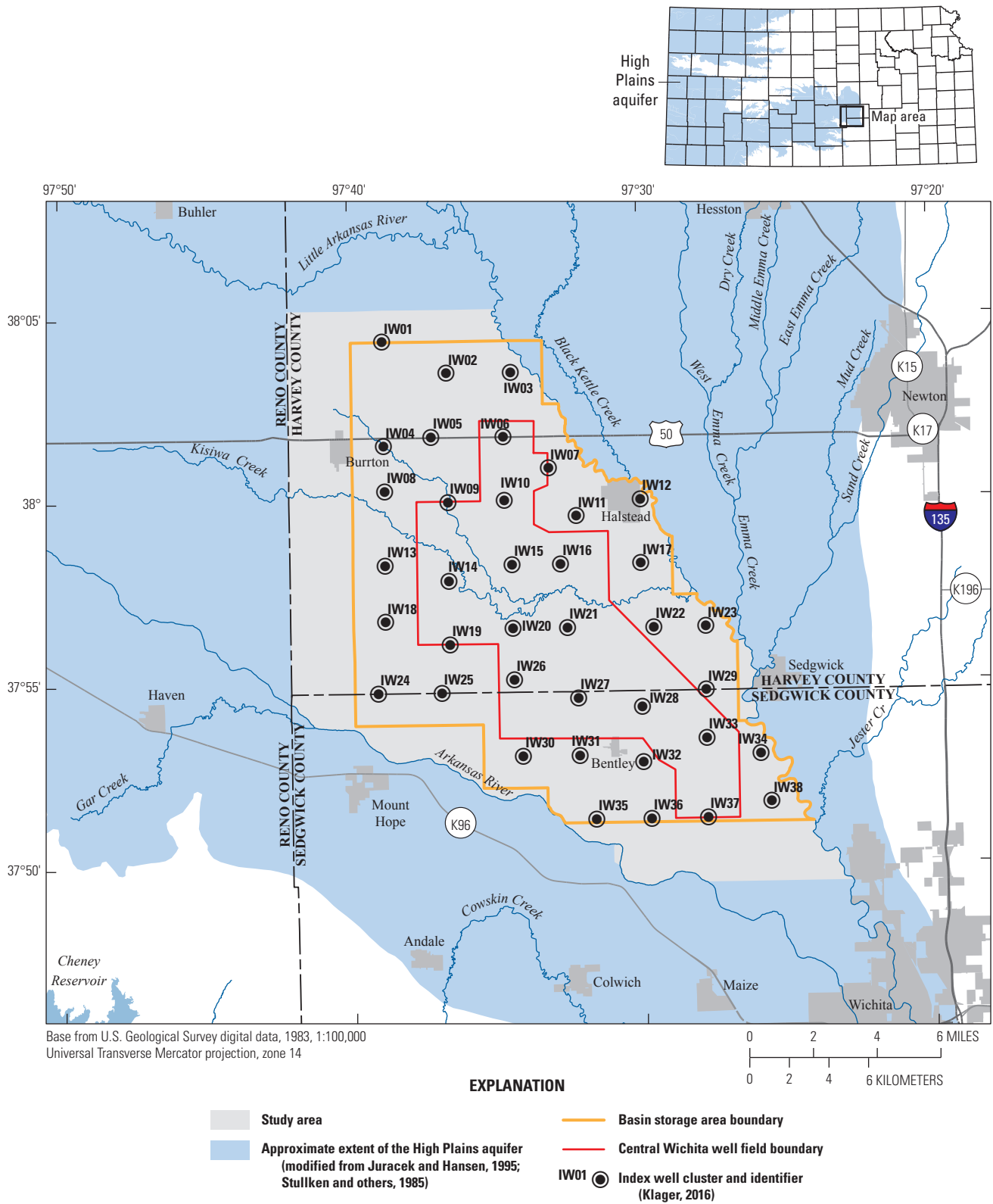


Figure 1. Location of the study area near Wichita, Kansas.

estimated using a groundwater-flow model (Kelly and others, 2013). The BSA is divided into 38 index cells and the groundwater-flow model is used to calculate the effects of ASR project recharge on the flow budget of each cell. Within each index cell there is a cluster of two index wells, one shallow and one deep, which are used for monitoring water quantity and quality in the aquifer. Credits accumulated during wet years can be used during periods of drought when the city's other primary water source, Cheney Reservoir, might not have the available storage to meet the city's normal demand (City of Wichita, [2007?]).

Purpose and Scope

The purpose of this report is to present the results of numerous groundwater-level measurements from wells screened in the *Equus* Beds aquifer near Wichita, Kans., in January 2016. Potentiometric surfaces were interpolated from the groundwater levels for the shallow and deep parts of the aquifer, and maps of the potentiometric surfaces are presented in this report. The potentiometric-surface maps were used to estimate storage-volume change in the study area between predevelopment (the time period before substantial pumpage began in the area, defined in this report as before 1940) and January 2016; the historic low in 1993 and January 2016; and January 2015 and January 2016. Potentiometric surfaces were used to generate maps showing the spatial distributions of estimated groundwater-level change between January 2015 and January 2016; predevelopment and 1993; and 1993 and January 2016 in the study area, which are also presented in this report.

Description of Study Area

The study area (fig. 1) is located in south-central Kansas, northwest of the city of Wichita. The study area is within the boundary of the *Equus* Beds aquifer, which is the easternmost part of the High Plains aquifer in Kansas. The study area is located between the Arkansas and Little Arkansas Rivers, overlying 189 square miles (mi²) of southwest Harvey and northwest Sedgwick Counties. The WWF is in the central part of the study area, with an area of about 55 mi². Most of the BSA is within the study area (the northeast corner of the BSA extends outside of the study area), with a total area of about 141 mi².

The study area has a continental climate characterized by cold winters and hot summers. In 2015, at a National Weather Service weather station at the Dwight D. Eisenhower National Airport in Wichita, the maximum recorded temperature was 102 degrees Fahrenheit (°F) on September 7, and the minimum temperature recorded was 0 °F on January 7 (National Oceanic and Atmospheric Administration, 2016a). In 2015, the average total precipitation across the study area was about 41.8 inches (National Oceanic and Atmospheric Administration, 2016b). The 30-year average

annual precipitation in the study area is about 32.7 inches, based on observed precipitation datasets from the National Weather Service Advanced Hydrologic Prediction Service (National Oceanic and Atmospheric Administration, 2016b).

In the study area, the *Equus* Beds aquifer consists mostly of Quaternary-age sand and gravel with some interbedded clay. There are dune sands in the northern part of the study area that form a topographic high (Williams and Lohman, 1949). The average aquifer thickness is estimated to be about 110 feet (ft), with the thickest part of the aquifer at about 330 ft (Kelly and others, 2013). Groundwater-flow direction is generally west to east, with the Arkansas River discharging into the aquifer in the west and the aquifer discharging into the Little Arkansas River in the east.

Land use is mostly crop production and grasslands. Based on data compiled by the Kansas Applied Remote Sensing Program (KARS) in 2005, land cover was about 73 percent cropland and about 20 percent grassland, with the remainder split between woodlands, urban lands, water, and conservation reserve program lands (Kansas Applied Remote Sensing Program, 2010).

Most groundwater use in the study area is for municipal or agricultural purposes. Outside of the WWF, the primary use is irrigation; inside the WWF, the primary use is municipal. In the study area, irrigation typically has an inverse relationship with precipitation, with higher withdrawals in years of lower precipitation (fig. 2). Before 1993, the municipal water use in the WWF was about 40,000 acre-feet per year (acre-ft/yr). With the adoption of the ILWSP, municipal groundwater use from the WWF decreased to about 20,000 acre-ft/yr by the late 1990s. The city of Wichita substantially decreased pumping from the WWF during 2014 and 2015 (fig. 2; Kansas Geological Survey and Kansas Department of Agriculture, 2016; Ginger Pugh, Kansas Department of Agriculture, written commun., 2016).

Previous Studies

Water levels and storage changes in the study area have been documented in previous years (Hansen, 2009, 2011, 2012; Hansen and Aucott, 2010; Hansen and others, 2013, 2014; Whisnant and others, 2015). Hansen and others (2013) mapped the 1993 historic low water levels. Kelly and others (2013) developed a groundwater-flow model of the *Equus* Beds aquifer for the purpose of estimating storage changes in the BSA attributable to the ASR project artificial recharge. Water quality in the study area has been documented by Ziegler and others (2010) and Tappa and others (2015).

Methods

This section describes the methods used to collect and process the groundwater-level data used in this report, interpolate potentiometric surfaces, and estimate changes in

4 Status of Groundwater Levels and Storage Volume in the *Equus* Beds Aquifer near Wichita, Kansas, January 2016

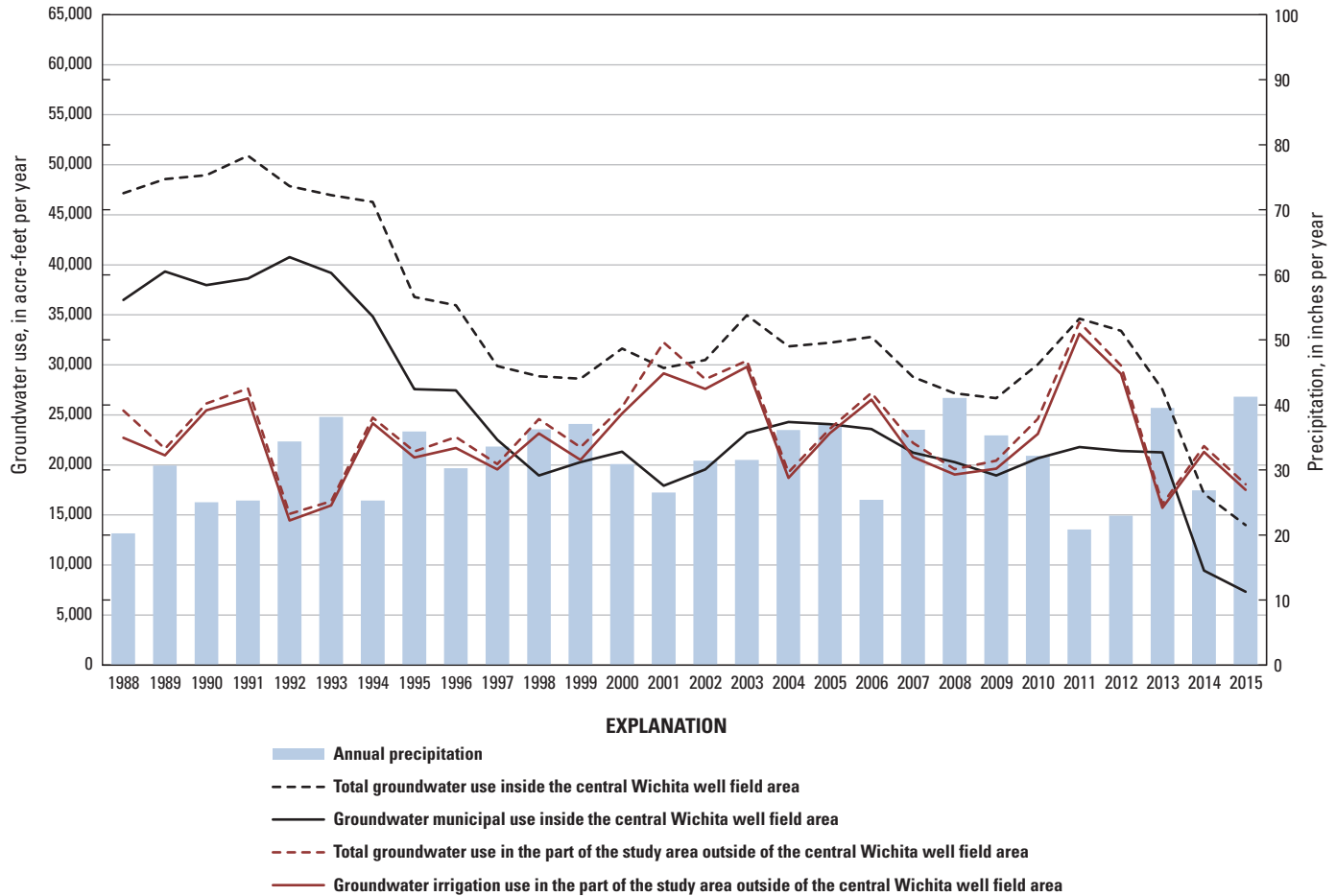


Figure 2. Annual groundwater use in the central Wichita well field area and the rest of the study area, and average annual precipitation in the study area, 1988 through 2015 (Water-use data are from Kansas Geological Survey and Kansas Department of Agriculture, 2016; precipitation data are from the National Oceanic and Atmospheric Administration, 2016a, 2016b) (modified from Hansen and others, 2014; Whisnant and others, 2015).

aquifer storage volume in the study area. These methods are similar to those described in previous reports (Whisnant and others, 2015; Hansen and others, 2014; Hansen and others, 2013).

Groundwater-Level Measurements

Groundwater levels were collected in numerous wells screened in the *Equus* Beds aquifer in and around the study area (fig. 1) from January 4 through January 29, 2016. The water levels were measured by the city of Wichita, the *Equus* Beds Groundwater Management District No. 2 (GMD2), the Kansas Department of Agriculture-Division of Water Resources (KDA-DWR), and the U.S. Geological Survey. Pumping from the WWF was paused in mid-December to allow the groundwater levels to recover to static or near-static conditions prior to the measurements. Groundwater-level measurements were completed following the procedures described in Cunningham and Schalk (2011). The groundwater data are stored in the U.S. Geological Survey's

National Water Information System (NWIS). These data are available to the public through the NWIS web interface (U.S. Geological Survey, 2016b). A total of 477 measurements collected in and around the study area during January 2016 were used for this report. The measurements used were selected based on whether the well depth was known and whether the well was needed to achieve adequate spatial coverage of the study area to perform an interpolation of the head values for the potentiometric-surface maps of the shallow and deep parts of the aquifer. The point measurements used to perform the analyses described in this report are compiled in a data release (Klager, 2016).

Shallow and Deep Parts of the Aquifer

The *Equus* Beds aquifer is thought to be unconfined in most of the study area; however, some shallow and deep wells that are clustered together have substantial differences in measured water levels, indicating confinement or semi-confinement of the deeper zone. Therefore, wells were

classified as being screened in the shallow or deep part of the aquifer based on two criteria. First, many wells have already been assigned to the shallow or deep layer of the aquifer and are named accordingly. For instance, each ASR index monitoring well cluster has a shallow and deep well, with well names ending in “A” or “C,” indicating that they are screened in shallow or deep parts of the aquifer, respectively. Second, if no indication is given in the well name, wells are classified into shallow and deep parts of the aquifer based on well depth. Ziegler and others (2010) classified wells less than 80 ft deep as shallow wells and wells 80 ft deep or deeper as deep wells, and this convention is used in this report.

Interpolation of Potentiometric Surfaces

The shallow and deep potentiometric surfaces were interpolated using the ArcGIS tool “Topo to Raster” (Esri, 2014) in a geographic information system (GIS). The initial surface interpolations used only the point groundwater-level measurements. Because the initial interpolations had some features that were unrealistic for the potentiometric surfaces, the initial surfaces were contoured, and the contours were manually adjusted based on hydrologic expertise. The Topo to Raster tool was used to reinterpolate the surfaces using the point groundwater-level measurement data and the manually adjusted contours to produce the final potentiometric surfaces. The final interpolated rasters and potentiometric contour files are available as part of a data release (Klager, 2016).

Estimation of Storage-Volume Change

Changes in saturated aquifer volume were estimated by calculating the groundwater-level change between the shallow January 2016 potentiometric surface and past shallow potentiometric surfaces in each raster cell in the GIS. Groundwater-level differences were then multiplied by the area of the raster cell to calculate the change in saturated aquifer volume in each cell. Changes in saturated aquifer volume in all cells were summed to calculate the change in saturated volume of the aquifer in the study area. Changes in storage volume were estimated for predevelopment to January 2016, the historic low of 1993 to January 2016, and January 2015 to January 2016. The January 2015 potentiometric surface was estimated by Whisnant and others (2015). The 1993 and predevelopment potentiometric surfaces were estimated by Hansen and others (2013).

To convert a change in saturated aquifer volume to a change in volume of water stored in the aquifer, the change in saturated volume must be multiplied by a specific yield value. Specific yield is the volume of water per unit bulk aquifer volume that will drain out under the force of gravity (Meinzer, 1923). Previous studies (Hansen and others, 2013, 2014; Whisnant and others, 2015) used a specific yield

value of 0.15, and that value was used in this report. This value was determined from calibration of a groundwater-flow model of the *Equus* Beds aquifer by Kelly and others (2013). Earlier reports (Hansen and Aucott, 2010; Hansen, 2009, 2011, 2012) had used a specific yield value of 0.20.

Potentiometric Surface Maps, January 2016

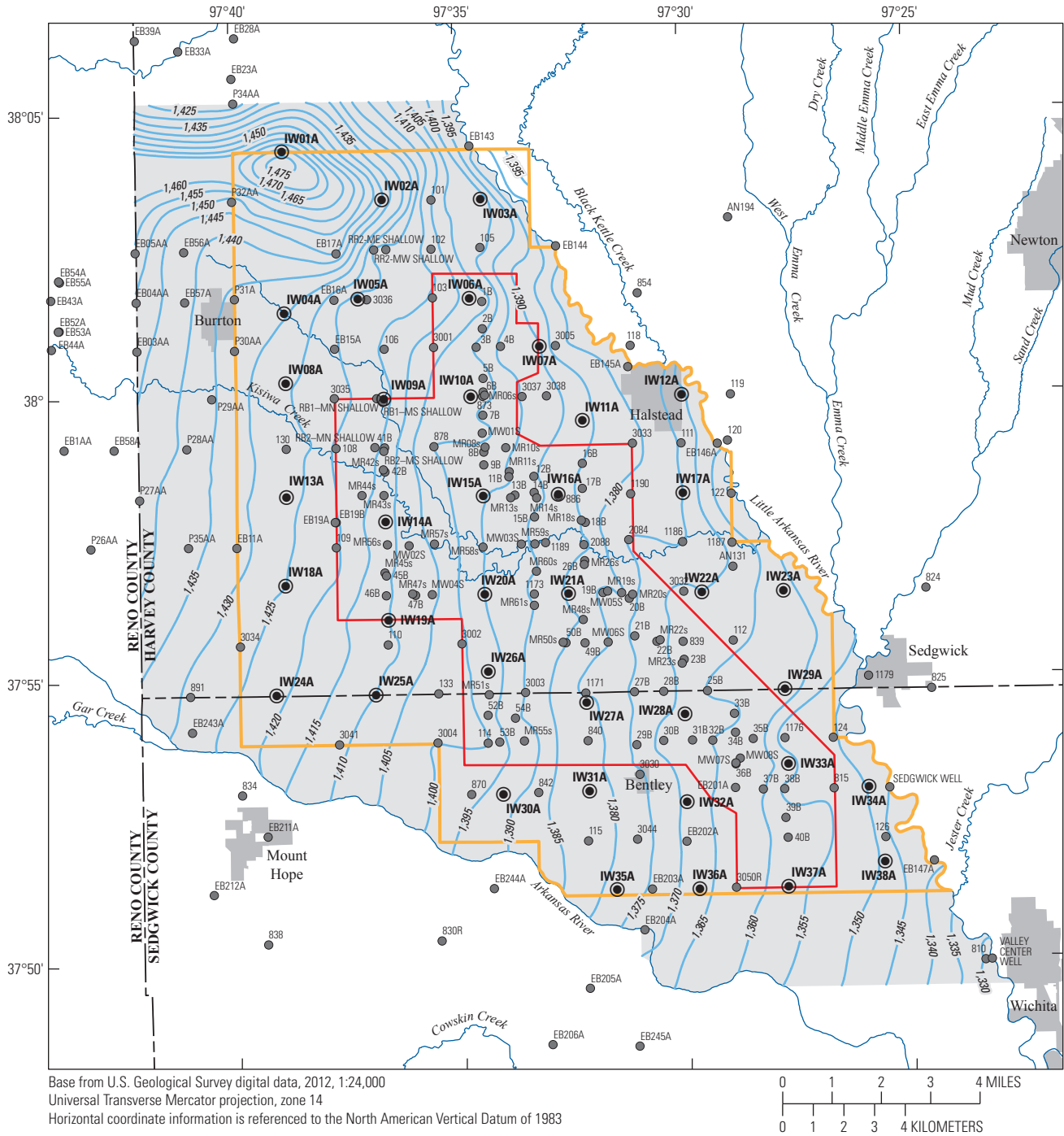
The potentiometric surfaces produced using the January 2016 groundwater-level measurement data are presented in figures 3 and 4. The groundwater-flow direction, as indicated by the head gradients in both shallow and deep maps, is west to east. The maximum head in the shallow surface was 1,475.89 ft, located in the northwestern part of the study area near monitoring well IW01A in the northwestern corner of the BSA. The minimum head value in the shallow surface was 1,325.95 ft in the southeastern corner of the study area near the Little Arkansas River. The maximum head value in the deep potentiometric surface was 1,443.93 ft at the western boundary of the study area south of Burton. The minimum head value in the deep potentiometric surface was 1,326.38 ft, located in the southeastern corner of the study area near the Little Arkansas River.

Groundwater-Level and Storage-Volume Changes in the *Equus* Beds Aquifer

Groundwater levels were generally higher in the study area in January 2016 than they were in January 2015 (fig. 5). Multiple factors contributed to increased levels, including increased precipitation in 2015 and less pumping by both agricultural and municipal users in 2015. The average groundwater-level increase in the shallow part of the aquifer between January 2015 and January 2016 was 3.37 ft in the study area, 3.82 ft in the BSA, and 4.50 ft in the WWF. These average groundwater-level changes were calculated by averaging the differences between the shallow potentiometric surfaces from January 2015 and January 2016 in the GIS. The groundwater-level changes in the shallow part of the aquifer are used to estimate changes in storage volume.

The volume of water stored in the study area, the BSA, and the WWF all increased from January 2015 to January 2016. The storage volume increased by about 61,000 acre-feet (acre-ft) in the study area, about 51,000 acre-ft in the BSA, and about 24,000 acre-ft in the WWF. The storage increase can most likely be attributed to lower volumes pumped by the city of Wichita (about 7,400 acre-ft in 2015 compared to about 9,700 acre-ft in 2014) (Ginger Pugh, Kansas Department of Agriculture, written commun., 2016)

6 Status of Groundwater Levels and Storage Volume in the *Equus* Beds Aquifer near Wichita, Kansas, January 2016



- EXPLANATION**
- Study area
 - Basin storage area boundary
 - Central Wichita well field boundary
 - 1,440 Potentiometric contour—Shows the altitude at which water would have stood in tightly cased wells in the shallow part of the *Equus* Beds aquifer, January 2016. Contour interval of 5 feet. Vertical Datum is North American Vertical Datum of 1988. Groundwater-flow direction is perpendicular to potentiometric contours from high to low values
 - EB10A Shallow monitoring well and identifier
 - IW01A Shallow index monitoring well and identifier (Klager, 2016)

Figure 3. Potentiometric surface of the shallow part of the *Equus* Beds aquifer study area, south-central Kansas, January 2016.

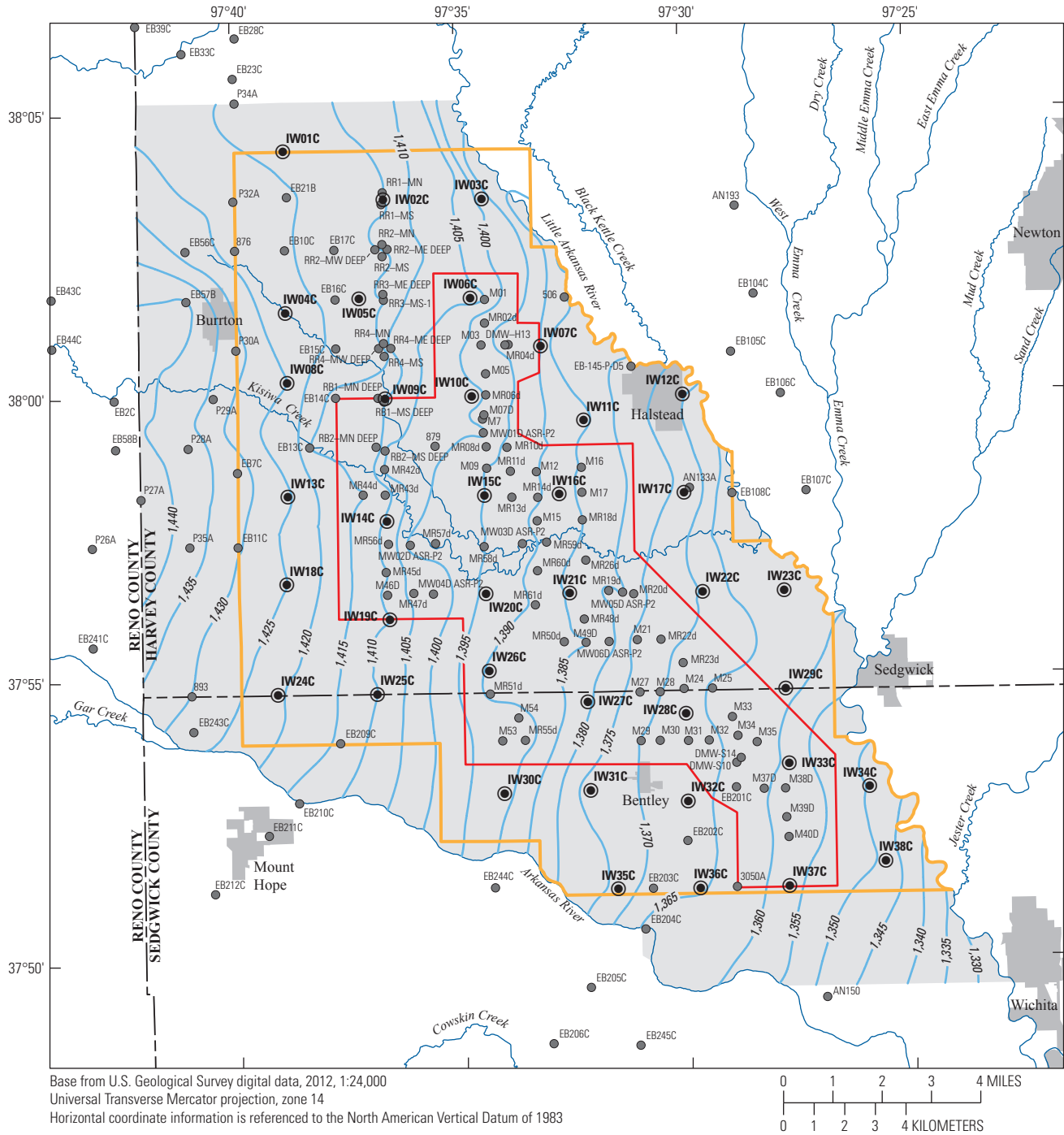
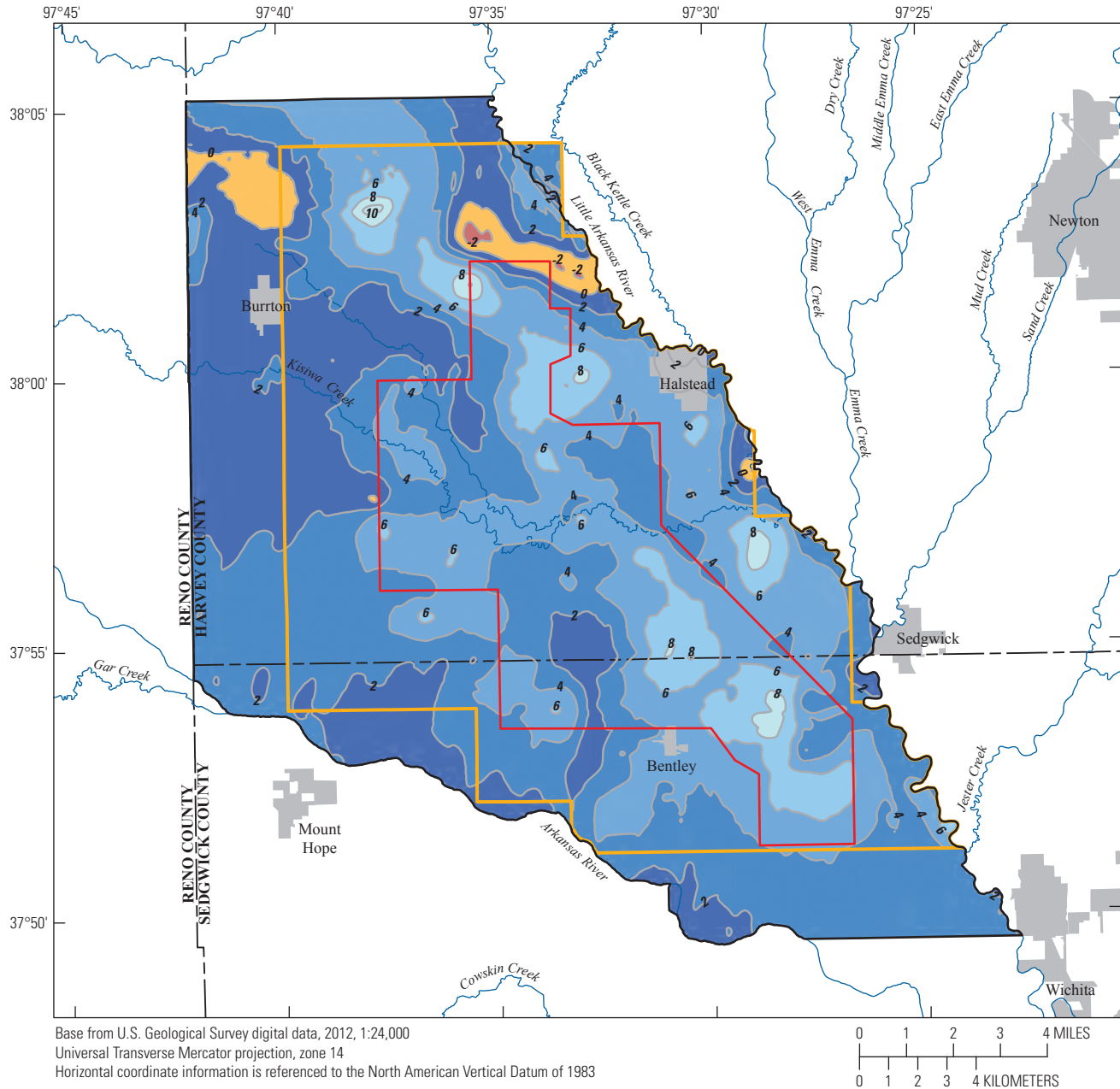


Figure 4. Potentiometric surface of the deep part of the *Equus* Beds aquifer study area, south-central Kansas, January 2016.

8 Status of Groundwater Levels and Storage Volume in the *Equus* Beds Aquifer near Wichita, Kansas, January 2016



EXPLANATION

Groundwater-level change in the shallow part of the *Equus* Beds aquifer, January 2015 to January 2016

- Decline of 2 to less than 4 feet
- Decline of zero to less than 2 feet
- Rise of zero to 2 feet
- Rise of greater than 2 to 4 feet
- Rise of greater than 4 to 6 feet
- Rise of greater than 6 to 8 feet
- Rise of greater than 8 to 10 feet

- Study area
- Basin storage area boundary
- Central Wichita well field boundary
- Contour of equal groundwater-level change, January 2015 to January 2016. Contour interval, 2 feet

Figure 5. Groundwater-level change in the shallow part of the *Equus* Beds aquifer study area in south-central Kansas, January 2015 to January 2016.

and higher precipitation leading to more recharge of the aquifer (about 41.3 inches of precipitation in 2015 compared to about 27.7 inches in 2014, with a 30-year average precipitation of 32.7 inches across the study area) (National Oceanic and Atmospheric Administration, 2016b). There was also an increased volume of artificially recharged water in 2015. The ASR project recharged about 1,770 acre-ft in

2015 compared to about 950 acre-ft in 2014 (U.S. Geological Survey, 2016a). The January 2016 storage volume in the study area was about 74,000 acre-ft less than the predevelopment storage volume. The storage-volume increase since the historic low of 1993 was about 121,000 acre-ft, a 62-percent recovery of the storage lost from predevelopment to 1993 (table 1, figs. 6–9).

Table 1. Storage-volume changes in the *Equus Beds* aquifer near Wichita, Kansas, since predevelopment (pre 1940) and since 1993 to January 2016 for the study area, the basin storage area, and the central Wichita well field area.

[--, not applicable]

End of time period	Storage-volume changes	
	Since predevelopment (acre-feet)	Since 1993 (acre-feet)
Study area		
1993	¹ -195,000	--
January 2012	¹ -125,000	² 70,000
July 2012	¹ -175,000	² 20,000
January 2013	¹ -154,000	² 41,000
January 2014	¹ -116,000	² 79,000
January 2015	¹ -135,000	¹ 60,000
January 2016	-74,000	121,000
Basin storage area		
1993	¹ -188,000	--
January 2012	¹ -112,000	² 76,000
July 2012	¹ -155,000	² 33,000
January 2013	¹ -134,000	² 54,000
January 2014	¹ -108,000	² 80,000
January 2015	¹ -119,000	¹ 69,000
January 2016	-68,000	120,000
Central Wichita well field area		
1993	¹ -121,000	--
January 2012	¹ -59,000	² 62,000
July 2012	¹ -74,000	² 47,000
January 2013	¹ -69,000	² 52,000
January 2014	¹ -63,000	² 58,000
January 2015	¹ -65,000	¹ 56,000
January 2016	-41,000	80,000

¹Storage-volume change previously reported in Whisnant and others (2015).

²Storage-volume change previously reported in Hansen and others (2014).

10 Status of Groundwater Levels and Storage Volume in the *Equus* Beds Aquifer near Wichita, Kansas, January 2016

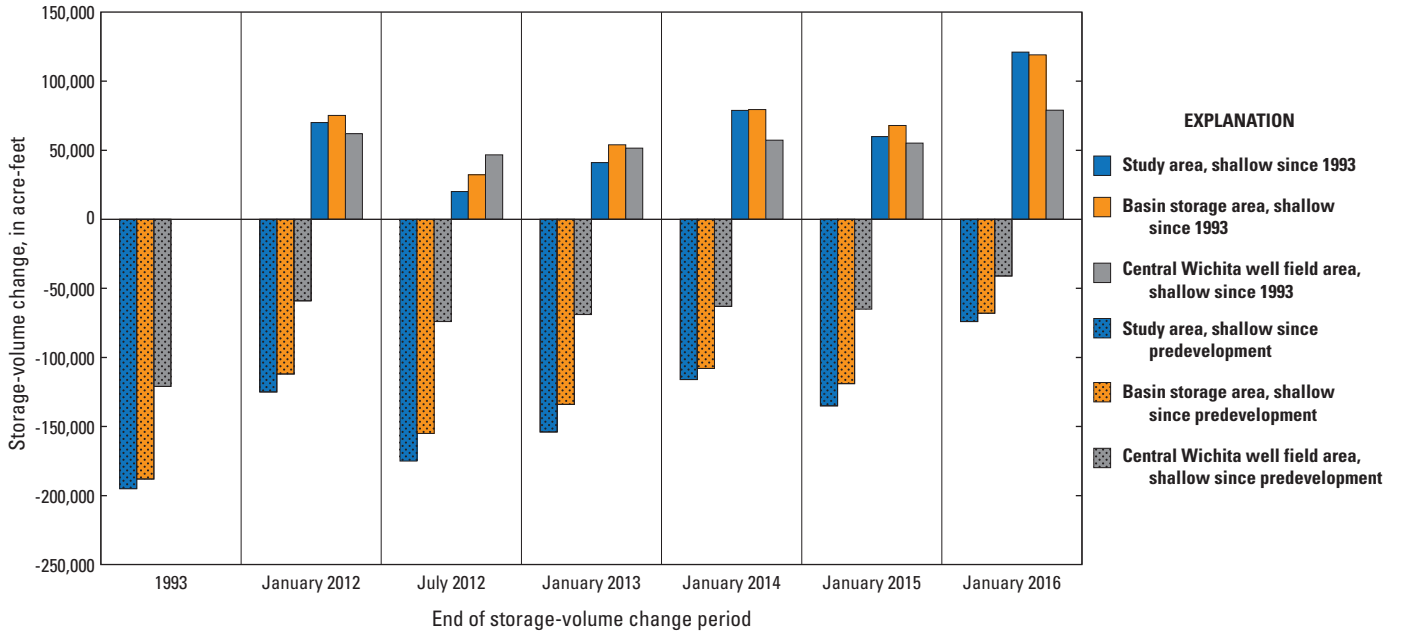
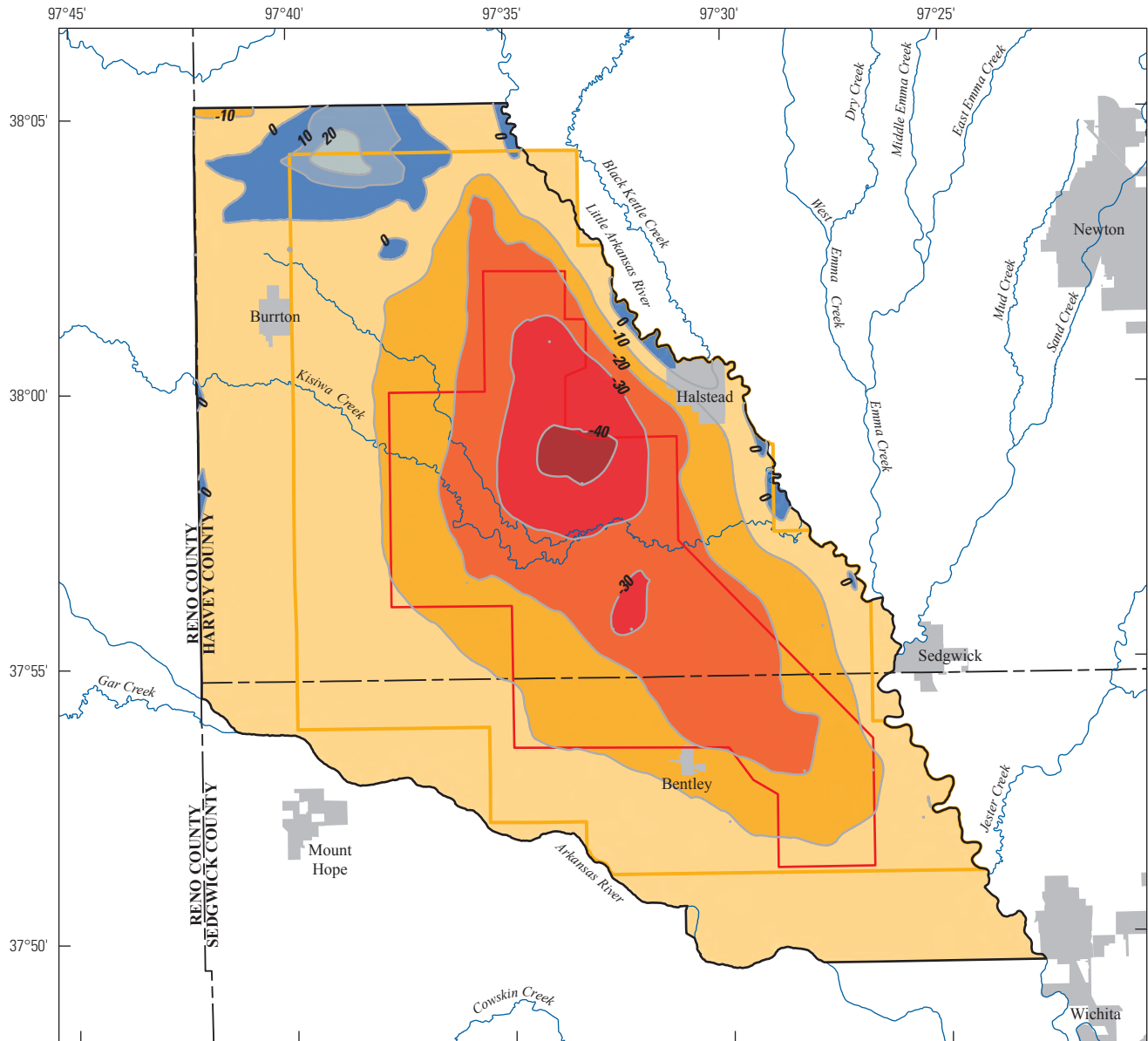
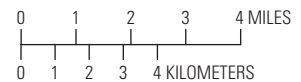


Figure 6. Estimates of storage-volume change since predevelopment (before 1940) and 1993 in the *Equus* Beds aquifer study area.



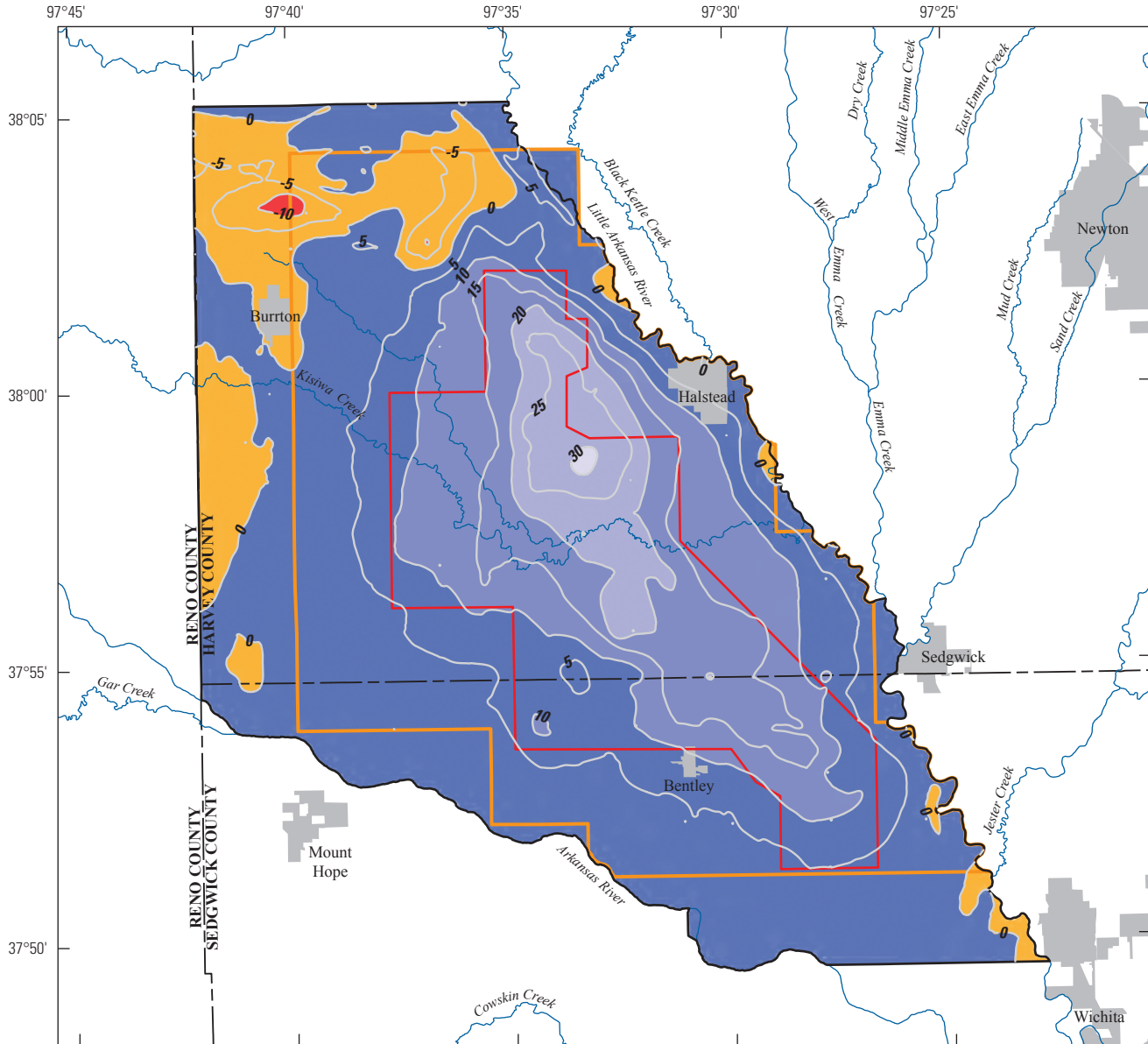
Base from U.S. Geological Survey digital data, 2012, 1:24,000
 Universal Transverse Mercator projection, zone 14
 Horizontal coordinate information is referenced to the North American Vertical Datum of 1983



EXPLANATION

- | | |
|---|---|
| Area of groundwater-level change | Study area |
| Decline of 40 feet or more | Basin storage area boundary |
| Decline of 30 to less than 40 feet | Central Wichita well field boundary |
| Decline of 20 to less than 30 feet | -20— Line of equal groundwater-level change in the shallow part of the <i>Equus</i> Beds aquifer between predevelopment and 1993. Contour interval, 10 feet |
| Decline of 10 to less than 20 feet | |
| Decline of zero to less than 10 feet | |
| Rise of zero to less than 10 feet | |
| Rise of 10 to less than 20 feet | |
| Rise of 20 feet or more | |

Figure 7. Groundwater-level changes in the *Equus* Beds aquifer in and around the Wichita well field in south-central Kansas between predevelopment (before 1940) and the historic low of 1993.



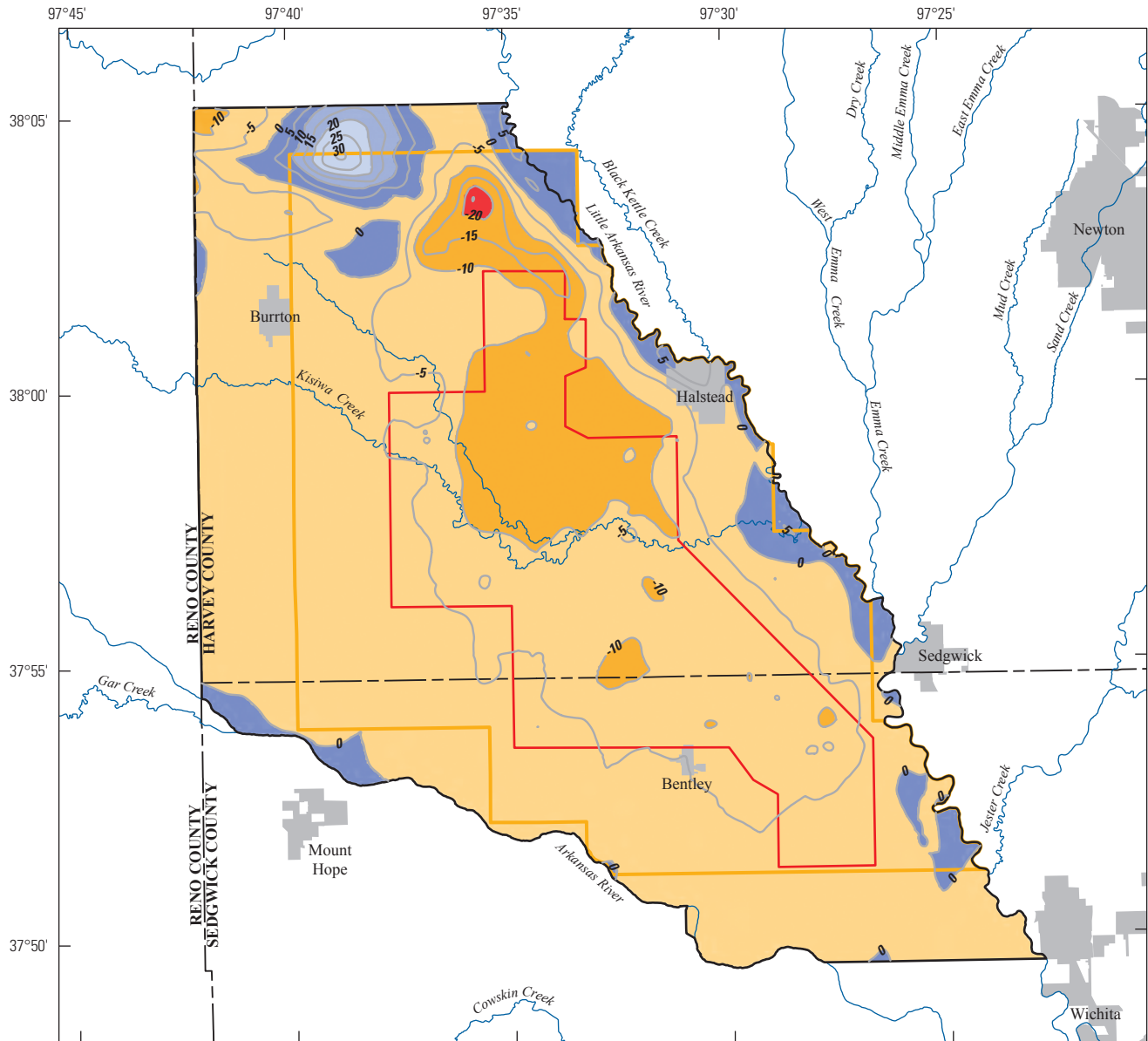
Base from U.S. Geological Survey digital data, 2012, 1:24,000
 Universal Transverse Mercator projection, zone 14
 Horizontal coordinate information is referenced to the North American Vertical Datum of 1983



EXPLANATION

- | | |
|---|--|
| Area of groundwater-level change | — Study area |
| Decline of 10 feet or more | — Basin storage area boundary |
| Decline of zero to less than 10 feet | — Central Wichita well field boundary |
| Rise of zero to less than 10 feet | —20— Line of equal groundwater-level change in the shallow part of the <i>Equus</i> Beds aquifer between 1993 and January 2016. Contour interval, 5 feet |
| Rise of 10 to less than 20 feet | |
| Rise of 20 to less than 30 feet | |
| Rise of 30 feet or more | |

Figure 8. Groundwater-level changes in the shallow part of the *Equus* Beds aquifer in and around the Wichita well field in south-central Kansas between 1993 and January 2016.



Base from U.S. Geological Survey digital data, 2012, 1:24,000
 Universal Transverse Mercator projection, zone 14
 Horizontal coordinate information is referenced to the North American Vertical Datum of 1983



EXPLANATION

Area of groundwater-level change:	Study area
 Decline of 20 feet or more	 Basin storage area boundary
 Decline of 10 to less than 20 feet	 Central Wichita well field boundary
 Decline of zero to less than 10 feet	 -20— Line of equal groundwater-level change in the shallow part of the <i>Equus</i> Beds aquifer between predevelopment and January 2016. Contour interval, 5 feet
 Rise of zero to less than 10 feet	
 Rise of 10 to less than 20 feet	
 Rise of 20 feet or more	

Figure 9. Groundwater-level changes in the shallow part of the *Equus* Beds aquifer in and around the Wichita well field in south-central Kansas between predevelopment (before 1940) and January 2016.

Summary

The *Equus* Beds aquifer in south-central Kansas is part of the High Plains aquifer that serves as a source of water for municipal and agricultural users in the area. The city of Wichita has used the *Equus* Beds aquifer as one of its primary water sources since the 1940s. The aquifer in and around Wichita's well field reached historically low water levels in 1993, prompting the city to adopt new water-use and conservation strategies to ensure that future water-supply needs would be met. Part of the plan was to initiate a managed aquifer recharge program called the *Equus* Beds Aquifer Storage and Recovery project. The goal of the managed aquifer recharge program is to artificially recharge the *Equus* Beds aquifer with treated water from the Little Arkansas River. As part of the *Equus* Beds Aquifer Storage and Recovery project, the city of Wichita and the U.S. Geological Survey have partnered in a long-term cooperative study to monitor and describe the quantity and quality of the water in the *Equus* Beds aquifer and the Little Arkansas River.

In January 2016, the city of Wichita, the *Equus* Beds Groundwater Management District No. 2, the Kansas Department of Agriculture—Division of Water Resources, and the U.S. Geological Survey collected numerous groundwater levels from wells screened in the *Equus* Beds aquifer in the area in and around Wichita's well field. The measurements were used to interpolate potentiometric surfaces for shallow and deep parts of the aquifer in the study area. These potentiometric surfaces were compared with potentiometric surfaces from previous years to estimate changes in storage volume in the study area.

Groundwater levels were generally higher in January 2016 than they were in January 2015. On average, groundwater levels in the shallow part of the aquifer were about 3.4 feet higher and groundwater levels in the deep part of the aquifer were about 3.8 feet higher in January 2016 than in January 2015. The volume of water stored in the study area decreased by about 74,000 acre-feet between predevelopment (before substantial pumpage began in the 1940s) and January 2016; increased by about 121,000 acre-feet between the historic low in 1993 and January 2016; and increased by about 61,000 acre-feet between January 2015 and January 2016. About 62 percent of the storage volume lost between predevelopment and 1993 has been recovered. The increase in storage volume from January 2015 to January 2016 can probably be attributed to less pumping by the city of Wichita and irrigators, more recharge due to higher-than-average precipitation, and higher volumes of artificial recharge in 2015.

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Publishing support provided by:
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